

Neural net aided detection of astronomical periodicities in geologic records

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Abstract

Astronomically controlled variations in the Earth's climate induce cyclic trends in the sedimentary process and record (Milankovitch periodicity). One of the main difficulties to be solved in order to choose among the registered periodicities is the conversion from the spatial (i.e. recurrent variations along the stratal sequences) to the temporal domains of the astronomically induced frequencies present in the rock record. We discuss here how this problem can be circumvented by teaching a neural net how to recognize periodicities in the signal. The application to two sequences of shallow water carbonate deposits from the Cretaceous of Southern Italy has shown this approach to be particularly effective, confirming the existence of Milankovitch-type periodicities in the records examined, where climate, sediments and biota concomitantly react to the variation in the solar constant induced by secular perturbations of the Earth's orbital elements.

Keywords: Milankovitch theory; paleoclimatology; Cretaceous; Southern Apennines

1. Introduction

Periodic changes in climate, induced by astronomically induced variations in the distribution of solar energy over the Earth, have been recognized as influencing the production of carbonate sediments, which are mostly formed by marine organisms and by their activity [1–4]. On the other hand, these periodic signals may be recorded in and have been extracted from a variety of other non-carbonate de-

posits originated in different sedimentary environments, from continental to deep marine [1,4].

The whole topic presents two incongruent aspects. The availability of detailed dynamic models of the Earth–Moon–Sun system, as well as of the Solar System as a whole, together with reliable algorithms, has led to very accurate predictions of the secular variations in the Earth's orbital parameters [5], while the signature left by these periodic modulations in the stratigraphic record generally appears "extremely imperfect, unreliable, noisy and poorly time calibrated" [3]. This lack of congruent effects in the sedimentary record is the outcome of a complex

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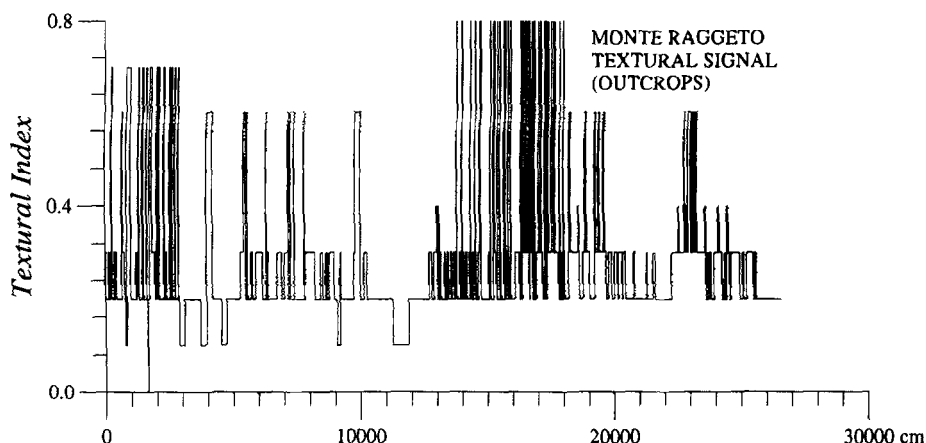


Fig. 1. Textural signal extracted from Mount Raggeto sequence. The horizontal scale gives the stratigraphic thickness of the sampled succession of strata in centimetres (younging to the right). Each interval in the vertical scale corresponds to a different standard texture (based on Embry and Klovan classification), as coded in Table 1.

process, driven by a number of variables, of which climate (together with sediment supply, global sea level change, subsidence, biologic activity, etc.) is just one. An extensive review of these processes can be found in a recently published volume of proceedings [3], as well as in other recent publications

[1,4,6]. In addition, we must consider that, if the time series analyzed are acquired with not very dense sampling, many of the high-frequency cycles may be lost in the rock record [3,4].

For these reasons high-frequency cyclicality has been more and more studied in modern stratigraphy

Table 1

Code used for the Mount Raggeto and Mount Tobenna deposits (see also Fig. 1 Fig. 2)

Mount Raggeto Textures	Mount Tobenna Lithofacies	Codes
Bindstone	"Stromatolitic" and "loferitic" boundstones (criptalgal laminites)	0
Mudstone	Aeolissaccus sp. and Thaumatoporella sp., wack and wack-packstone, locally with burrows	0.1
Wackestone	Large Miliolid pack and pack-wack with Thaumatoporella sp. and Ostracods	0.2
Packstone	Wack, wack-pack and pack thin levels with criptalgal laminites	0.3
Grainstone	Miliolid wack and wack-pack with benthonic foraminifers, Thaumatoporella sp., Aelisaccus sp., Dycyclina sp., Ostracods and micritized grains	0.4
Rudstone	Oncoid pack, pack-wack and grains with Dycyclina sp., benthonic foraminifers and small planktonics	0.5
Floatstone	Foraminifers wack-pack and pack with small planktonics	0.6
Undetermined	Foraminifer wack and mud-wack, with small planktonics and with bioturbations	0.7
Clay levels	Radiolitic floatstone with bioclastic matrix	0.8

Note that while for M. Raggeto only the textures (based on Embry and Klovan, 1971, standard classification) are used, a more detailed study has allowed determination of the lithofacies for M. Tobenna.

for its potential predictive value, in order to tie sedimentary processes to absolute time. In spite of the pioneer work of Croll and Gilbert [7,8] at the turn of the last century, it has been only in the past 15 years that an increasing interest has developed and the new term of 'cyclostratigraphy' has become established, stressing the genetic connection of the orbital perturbations induced by the variability of Earth's thermal budget (Milankovitch cyclicality) with the climate, sedimentation and biosphere [2]. The vast majority of works in cyclostratigraphy have been devoted to the pelagic realm; that is, where the biostratigraphy gives the best time definition, allowing for the detected cyclicities to be framed into the Milankovitch periodicity [2–4]. More recently, a number of authors have focused their attention on shallow marine carbonates, which seem to offer the best conditions for the detection of the shortest period components (see, e.g., [9–13]) even though, for these types of data, the chance of gaps punctuating the record may be high (e.g., in the peritidal domain).

In the Cretaceous sequences of the Southern Apennines more than 700 m of well bedded carbonate platform deposits, spanning from the Hauterivian to the Coniacian, have been analyzed on a centimetric scale. In these shallow water sequences three main rock types alternate and the relative lithofacies associations qualify them as subtidal, tidal–supratidal and storm-controlled deposits [9–12,14] (Table 1, Figs. 1 and 2). The cycles are aggradationally

stacked, shallow upwards and are often capped by ephemeral, emersion related features. Environmental changes, expressed by texture or lithofacies variations, plotted at a centimetre scale on the aggradation axis show a hierarchy of at least 3 orders of cycles [15].

A package of spectral analysis, based on analytical techniques originally tailored for astronomical applications, was used to search for possible periodicities [16]. This package includes pre-processing and spectral analysis of the signal; it deals with unevenly spaced data; it allows discrimination between meaningful frequencies and spurious ones on a physical basis and, finally, reduces the effects of aliasing bias in the data.

The duration of cycles was calculated for about 420 m of well bedded sequences by comparing the Relative Ratio Sets (RRS) of the recurrence of sedimentary features (expressed in centimetres) with the RRS of orbital parameters (Precession, Obliquity and Eccentricity) calculated by Berger [17,5] for the Cretaceous (expressed in years). The two ratio sets show, for every sequence analyzed, a very good linear correlation ($r > 0.99$), suggesting that Cretaceous carbonate platform strata of Southern Italy have an allocyclic organization (their cyclicality is not inherent to the sediment deposition [11,12,14]), and that the relative time scale falls in the Milankovitch periodicities (Table 2). Furthermore, spectral analysis of palaeomagnetic data from a $\cong 90$ m thick

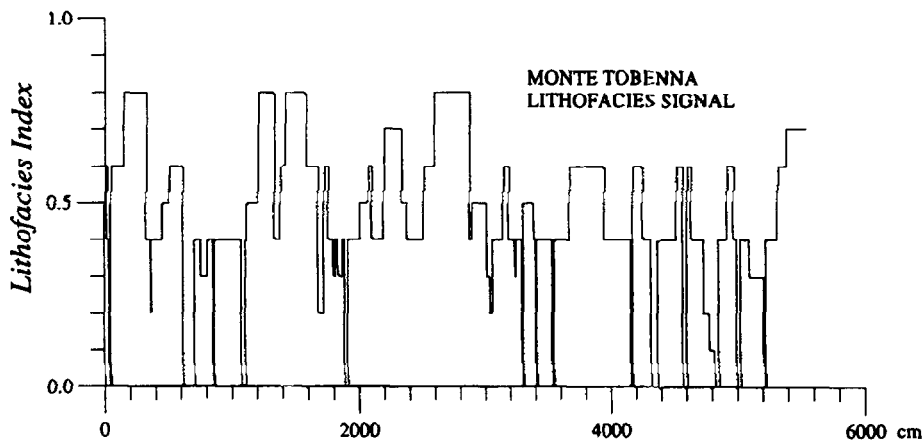


Fig. 2. Lithofacies signal extracted from Mount Tobenna sequence. On the horizontal scale the stratigraphic thickness is indicated (expressed in centimetres and younging to the right), while the vertical scale is punctuated with the 10 different lithofacies, reported in Table 1, recognized from a sedimentologic study at the outcrop as well as by using thin section and acetate peels.

Table 2
Space and time periodicities of shallow water carbonate deposits of Mount Tobenna (Upper Cretaceous) and Mount Raggeto (Lower Cretaceous)

TYPE OF SIGNALS	AGE	THICKNESS OF THE SEQUENCES	SIGNAL PERIOD. (cm)	SIGNAL PERIOD. RATIOS	CORR. COEFF.	ORBIT. RATIOS	ORBITAL PERIOD. (years)			AVERAGE ACCUM. RATE (CM/KY)		
							Eccentricity	Obliquity	Precession			
MONTE TOBENNA												
LITHOFACIES	Coniac.	60 m	1108	10.355	0.999	10.349	403600	95900	39000	2.75		
			264	2.467		2.459						
	Turon. pp		107	1.000		1.000						
MONTE RAGGETO												
TEXTURES (OUTCROP)	Albian	270 m	985	10.944	0.9995	10.568	403800	95800	49370	2.60		
				256		2.844					2.510	
				145		1.611					1.292	
			Barrem.	90		1.000					1.000	
BED THICKNESS (OUTCROP)	Albian	270 m	1773	10.707	0.9997	800000 ?	403800	95800	49370	2.47		
						985					2.489	10.568
						229					1.413	2.510
						130					1.000	1.292
			Barrem.			92					1.000	1.000
TEXTURES (CORE)	Barrem.	60 m	1947	22.902	0.9998	800000 ?	404200	94900	47650	2.92		
						1168					5.216	22.200
						266					2.725	5.211
						139					2.333	2.617
						119					1.274	2.084
						65					1.000	1.210
						51					1.000	1.000
			Haut. pp			258					4.962	5.211
PALBOMAGNETIC INTENSITY (CORE)	Barrem.	90 m	144	2.769	0.9994	94900	47650	37950	22000	2.85		
						103					1.981	2.084
						65					1.250	1.210
						52					1.000	1.000
			Haut. pp			1073					19.870	22.200
PALBOMAGNETIC INCLINATION (CORE)	Barrem.	90 m	260	4.815	0.9999	94900	47650	37950	22000	2.93		
						142					2.630	5.211
						118					2.185	2.617
						69					1.278	2.084
			Haut. pp			54					1.000	1.210

For the analytical data reported here see text and Fig. 5 and Fig. 6, Table 1 and [11,14,15,19,28]. Note that the accumulation rates are given here as an average.

core, drilled in lower Cretaceous shallow marine carbonate deposits from Mount Raggeto in Southern Italy (and partly overlapping with one of the sequences discussed in the present paper), also demonstrates cyclic recurrence of the remanent magnetization values (declination, inclination and intensity) [18,19]. These paleomagnetic periodicities are, again, linked to the orbital parameters with $r > 0.99$ (Table 2).

We discuss here a new approach based on neural networks, to identify meaningful periodicities, a tool which has proved particularly effective in speech processing [20], and in pattern recognition [21]. In Section 2, the data are presented, together with a short summary of the classical ‘power spectrum’ approach to the search for periodicities in stratigraphic records. The general background of the neural network implementation and training is introduced in Section 3, data reduction and analysis are detailed in Section 4, while results and concluding remarks are presented in Section 5. Finally, in Appendix A we provide the reader with some details on the training of the neural network.

2. The data

The data used for the present analysis refer to two stratigraphic sequences sampled in a few quarries at Mount Raggeto ([15,11] and Mount Tobenna [14],

located near Caserta and Salerno, Southern Italy, respectively. The geological characterization of these sequences has been discussed at length elsewhere [15,14] and we refer to these papers for further details. We want, however, to stress a few points:

1. the data were obtained by identifying/sampling the rocks directly from the outcrop at a centimetric scale and the lithofacies were determined by supplementary examination using thin sections and acetate peels;
2. both sequences refer to carbonate strata formed in shallow water environments, within the photic zone, at a depth never exceeding — on average — a few or a few tens of metres and, therefore, in environments very sensitive to global sea level oscillations (eustasy);
3. the sedimentation rates in these environments are usually fairly high and, therefore, sampling at a centimetre scale provides a good time resolution.

3. Methodology

Neural nets are powerful tools in dealing with:

1. function approximation: it has been shown that they are universal function approximators [22,23];
2. classification and clustering: a neural net can learn from examples how to classify input patterns in a supervised ([24,25]) or unsupervised ([24,26,27]) manner.

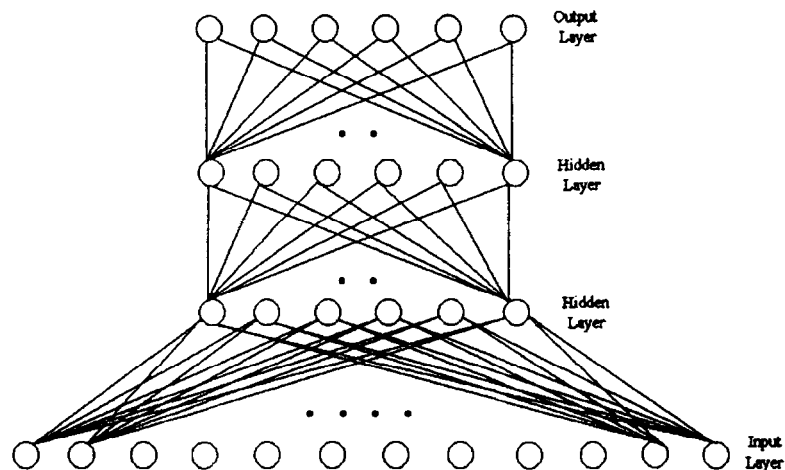


Fig. 3. The topological structure of the Multilayer Perceptron Neural Network used for our experiment.

In particular, their use is recommended when:

1. there is no good algorithm for solving the problem;
2. input data are incomplete, noisy and not directly understandable;
3. applications are data-intensive; that is, there are more data than computations;
4. the procedure needs to be repeated many times;
5. many examples for ‘training’ are available.

All these requirements are fulfilled by our data, thus making neural nets an ideal tool to deal with periodicity recognition in stratigraphic signals.

3.1. The neural net

A neural net is a computational structure made by many processing elements (units) — the neurons — operating in parallel. These neurons are generally organized into clusters or layers. They are grouped in ‘input’, ‘output’ and ‘hidden’ (i.e., those units which are neither input nor output) layers. Three fundamental elements characterize any neural nets:

1. the net structure or topology; that is, the way the layers are linked;
2. the activation function, which represents the answer of a neuron to the input stimuli;
3. the learning algorithm.

In the present case, the net model adopted is the well known ‘Multilayer Perceptron’ [25] shown in Fig. 3. It consists of one input layer, one output layer and one or two hidden layers. Each neuron in a given layer is connected to all the neurons of the next one.

Our model is synchronous: at each time every neuron receives as input the weighted sum of the

input patterns and/or of the other neuron outputs, as shown in the following equation:

$$O_k = f \left[\sum_n W_{kn} O_n - B_k \right] \quad (1)$$

where: W_{kn} is the weight associated to the link from neuron n to neuron k ; O_n is the output of neuron n or of the n th input; and B_k is the neuron threshold, generally called bias. Thus, the neuron output is a continuous and derivable function of its net input, with values comprised in the $[0,1]$ range. This function, f , is not linear for the hidden units. For our experiment, we chose the sigmoidal function, shown in Fig. 4, which looks like:

$$f(x) = 1 / (1 + e^{-x}) \quad (2)$$

The training procedure was the so-called ‘back propagation’, which works as follows: the first pattern is presented to the input neurons and then the net gives its output. If it is not equal to the desired output pattern, we compute the difference (error) between these two values and change the weights in order to minimize it. Then we propagate the information to the previous layers, changing the weights. For a detailed explanation of the algorithm see Appendix A. We repeat these operations for each input pattern until we minimize the mean square error of the system. Given the p th pattern in input, the error E_p is:

$$E_p = 1/2 \times \sum_j (t_{pj} - O_{pj})^2 \quad (3)$$

where t_{pj} is the p th desired output value of neuron j and O_{pj} is the output of the corresponding neuron.

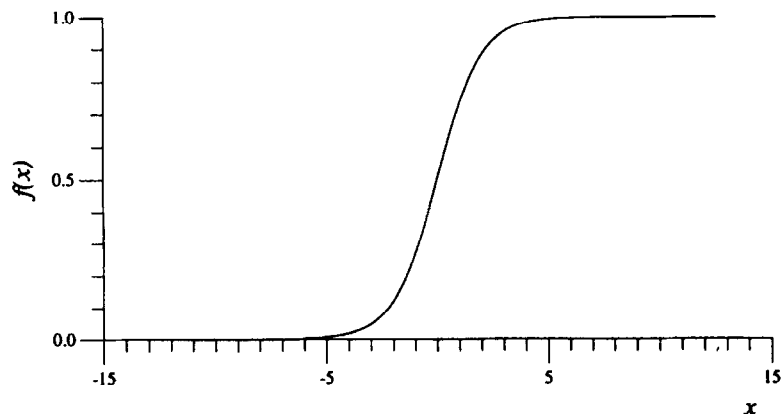


Fig. 4. Sigmoidal function used for the experiment of neural net training.

3.2. Application to geologic record

In this paper we address the specific problem of how to train a neural net to recognize the existence of those periodic signals which have been known for a long time to exist in biologic [26] as well as in stratigraphic signals [1,3,4]. Several authors have suggested that these signals might be linked to the secular variations of the Earth's orbit [5,3,17]. The main problem to be solved remains, however, how to convert the spatial frequencies observed in the signal to the time frequencies predicted by the orbital theory. Starting from the quite accurate time frequencies predicted for the Mesozoic, by Berger and his collaborators on the basis of N-body simulations of the solar system [5,17], we trained the neural network described above to recognize the existence of such periodic signals (regardless of their amplitude) in real data, once a rough estimate of the sedimentation rate S_{est} has been provided. This was achieved by producing a synthetic data string, based on S_{est} , and the set of frequencies estimated by Berger as the most likely for the assumed age of the strata considered. These strings were then used to train the net. In the following section we shall give more details on the application of the method to two specific cases.

4. Data analysis

Data analysis consisted of two main steps: (1) pre-processing aimed to reduce the noise level in the data; and (2) spectral analysis needed to have a first guess on the spatial harmonics present in the signal. The first goal was achieved by means of a simple running mean algorithm. The use of more refined filters would have been inappropriate due to the typology of the signal. The second step makes use, instead, of the so called Modified Scargle Algorithm [11,12] and leads to a rough estimate of S_{est} .

The Scargle algorithm has exactly the characteristics required. To be more precise:

1. it deals with unevenly spaced data; this characteristic is important because the rebinning of unevenly sampled data, like the stratigraphic signals, into equally spaced bins and the following computation of a conventional periodogram may alter the spectrum and the significance of a periodic signal;
2. it allows discrimination between meaningful frequencies and spurious ones on a physical basis; this is made by evaluating the false alarm probability [16]; that is, an estimate of the significance of the height of a peak in the power spectrum;

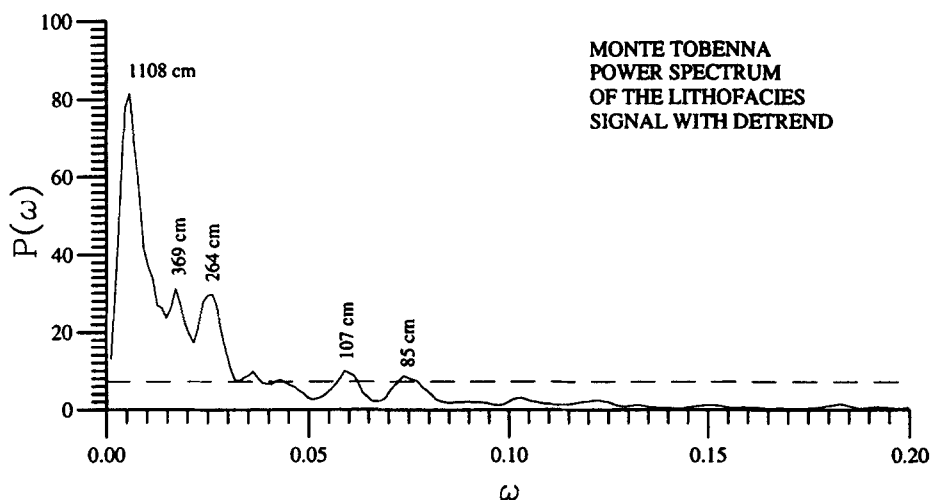


Fig. 5. Periodogram of the Mount Tobenna lithofacies. The signal has been tapered, to adjust spectral leakage, with a Gaussian window function. The horizontal dashed line gives the 90% confidence level.

this is possible only if the periodogram has an exponential probability distribution function [16], achieved by normalizing the periodogram with respect to σ^2 ; that is, the variance of the data;

3. it reduces the effects of aliasing bias in the data.

To be more detailed, let $X(t_i)$, with $i = 1, 2, \dots, N_0$, be the time series data, sampled at discrete intervals, and $P_X(\omega)$ the corresponding power spectrum at frequency ω . The equation to calculate $P_X(\omega)$ is:

$$P_X(\omega) = \frac{1}{2} \left[\frac{[\sum_j X_j \cos \omega(t_j - \tau)]^2}{\sum_j \cos^2 \omega(t_j - \tau)} + \frac{[\sum_j X_j \sin \omega(t_j - \tau)]^2}{\sum_j \sin^2 \omega(t_j - \tau)} \right] \quad (4)$$

where τ is defined as follows:

$$\tan(2\omega\tau) = \frac{(\sum_j \sin 2\omega t_j)}{(\sum_j \cos 2\omega t_j)} \quad (5)$$

and represents a term that makes the periodogram invariant to a shift in the zero point of the time scale.

4.1. Description of the experiments: Mount Tobenna and Mount Raggeto data

The stratigraphic record obtained for the exposed sequences in Mount Tobenna and Mount Raggeto

was processed according to the previous section, leading to the identification of the spatial periodicities shown in Figs. 5 and 6, and listed in Table 2.

4.1.1. Pre-processing

This is the most delicate step of any experiment done with neural networks and on it depend the choices of the training and test sets, as well as the best structure for the input data. The aim of our experiment was to teach the network how to recognize the possible existence of periodic signals in a very noisy spatial record, choosing between six possible classes; namely the six main periodicities expected for Milankovitch-type phenomena. In order to train the network we produced a series of simulated records having the following characteristics:

1. all possible combinations of the six Milankovitch periodicities;
2. additive noise;
3. the same length of the real stratigraphic sequence;
4. a similar square-like shape;
5. the same sampling rate as the real data set.

Due to the different domains of the real (space) and simulated (time) series, it was first necessary to find the conversion factor (sedimentation rate) α . In order to estimate α we proceeded as follows: from a preliminary spectral analysis of the data we derived the most significant peak and assumed it to be related to the highest Milankovitch frequency, as in [5,17], thus obtaining an estimate of α . This value

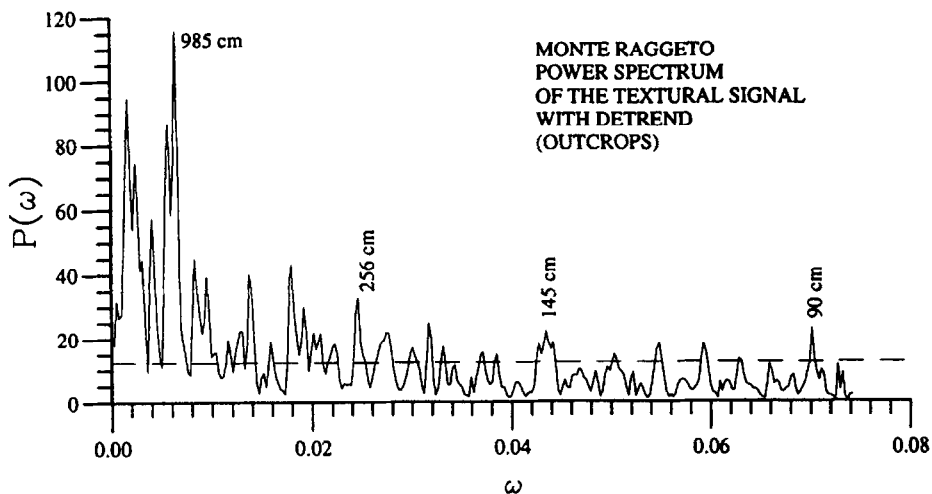


Fig. 6. Periodogram of the Mount Raggeto texture. The signal has been tapered, to adjust spectral leakage, with a Gaussian window function. The horizontal dashed line gives the 90% confidence level.

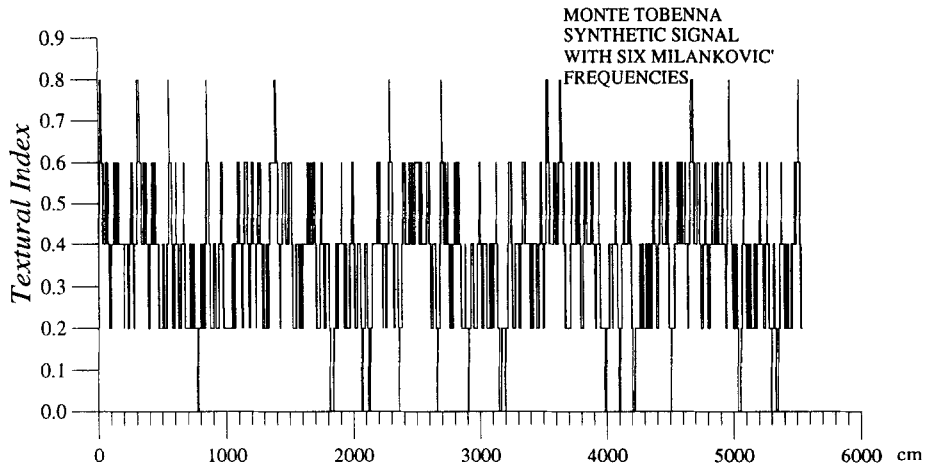


Fig. 7. Synthetic signal obtained for the Mount Tobenna texture record by combining the six Milankovitch cycles and assuming $S_{est} = 3.71 \text{ cm} \times \text{kyz}^{-1}$.

gave the sampling rate for the synthetic signals, which were therefore degraded by adding randomly generated additive noise. Fig. 7 shows an example of synthetic signal, containing all six Milankovitch frequencies, created by our procedure.

4.1.2. Creation of training and test sets

The net model used for these experiments is not shift-invariant and it is also impossible to submit the whole signal to the net at the same time. In order to by-pass this problem, we subdivided every simulated data set into smaller subsequences long enough to contain all the information. Every subsequence is slightly overlapping the next in order to simulate a sort of continuity in submitting the signals to the net. For each subsequence, a Fourier transform allows identification of the six amplitude values which represent the pattern input; so every signal is submitted

to the network as a pattern set. We give the net two consecutive overlapped sequences in order to obtain the best experimental results. In this way we obtained the 12 input values of the single input pattern. Each time series is composed of 10 and 13 consecutive overlapped sequences for the Mount Tobenna and Mount Raggeto experiments, respectively (depending on the length of the related original stratigraphic signal); we obtained 9 and 12 input patterns for each synthetic signal related to the Tobenna and Raggeto experiments respectively. We have 64 (2^6) different synthetic signals, containing all the combinations of the 6 Milankovitch periodicities [(frEq. (1)), (frEq. (2)), ..., (frEq. (6)), (frEq. (1) and frEq. (2)), ..., (frEq. (1) and frEq. (6)), (frEq. (2) and frEq. (3)), ..., (frEq. (2) and frEq. (6)), ..., (frEq. (1) and frEq. (2) and frEq. (3) and frEq. (4) and frEq. (5) and frEq. (6))] plus an additive pseudo-random noise.

Table 3
The training phase

Synthetic signal	Network topology	Cycles	RMS - Trains
Tobenna	12-6-6-6	77,000	0.013117
Raggeto	12-6-6-6	95,000	0.009991

The network topology contains the information on the number of layers and the number of neurons per layer, respectively: "in-hid1-hid2-out".

Table 4
The test phase

Signal	RMS - Test	Target periodicities	Recognized periodicities
Tobenna	0.124958	F1,F2,F3,F4,F5	F1,F2,F3,F4,F5
Raggeto	0.031333	F1,F2,F3,F4	F1,F2,F3,F4

F_i , $i = 1, \dots, 6$ are the six Milankovitch periodicities, from the highest ($F_1 \cong 400$ Ky) to the lowest ($F_6 \cong 18$ Ky).

The training set is composed of 30 synthetic signals, chosen from the 64 available, in such a way that samples of the 6 fundamental periodicities and of their combinations are present. For each time series the output error is evaluated as the mean value of the output of the 9 or 12 consecutive input patterns. The error is back-propagated after each submission of all 30 synthetic signals to the net.

4.2. Experimental results

We used the pre-processing strategies, described in Section 4.1.1, to evaluate the sedimentation rate and the sampling rate for the sequences related to Mount Tobenna and Mount Raggeto. Once the sampling rate was fixed, a complete series of synthetic signals covering all possible combinations of the six main Milankovitch frequencies was produced. We then applied the procedure described in Section 4.1.2 to create the training and test set. After the training phase (see also Appendix A we used a subset of synthetic signals not used in the training phase in order to evaluate the network performance. With these signals we obtained results with 100% of correct output detection (with an output value greater than 0.9 for the output nodes related to the Mi-

lankovitch cycles present into the signals). Finally, we used the Tobenna and Raggeto sequences in order to evaluate the network behaviour with real signals. The results are illustrated in Table 3, Tables 4 and 5. We used a threshold on the output equal to 0.5 for the detection. Therefore, this procedure allowed us to recognize the Milankovitch periodicities present in the stratigraphic signals studied. These results are confirmed by spectral analysis developed with the Scargle algorithm and the evidence from the outcrop (thickness of the cycle) [14,11].

The neural net approach here illustrated is completely automatic and permits the users to detect Milankovitch periodicities without complex interpretations of the power spectrum.

5. Conclusions

We have presented a neural net approach to the recognition of periodicities in the stratigraphic record. With respect to the traditional techniques based on the power spectrum (PST), this method offers several methodological and practical advantages which render the detection of periodicities much more reliable. In fact, instead of searching blindly for peaks in the

Table 5
The test phase

Signal	Freq. 1	Freq. 2	Freq. 3	Freq. 4	Freq. 5	Freq. 6
Tobenna	99.48	97.77	87.69	63.51	57.44	0.38
Raggeto	82.17	85.96	55.18	57.84	41.22	16.53

The output value of the output nodes corresponding to the 6 Milankovitch frequencies. The values are expressed in centesimals. Since the threshold is 50.0, we recognized the frequencies given in Table 4.

periodogram, the neural net approach presented here makes use of PST only to obtain a first-order estimate of the sedimentation rate. The existence of periodic components in the stratigraphic record is then performed on a yes/no base by training the net to recognize signals of a given frequency. The time and effort required for training of the net is largely compensated for by the effectiveness of the method and by the relatively short computing time required by the following processing steps.

The method has been tested on Cretaceous sequences, formed by well bedded carbonate rocks of shallow water origin. The data processed were obtained with centimetric accuracy along two exposed sections at Mount Raggeto and Mount Tobenna, Southern Italy, which had been studied previously by the authors with traditional PST's [11,14]. The texture and lithofacies based analysis had already showed a repetitive organization of the sedimentary environments, suggesting cyclical alternation of deeper and shallower depositional settings up to emersions [14,15]; this behaviour was assumed to be controlled by Earth's orbital perturbation [11]. The results of the present study confirm the existence of Milankovitch type signals ($\cong 20\text{--}\cong 400\text{ Ky}$) in both sequences and, in particular:

1. Eccentricity (long and short), Obliquity (long and short) and Precession (long) terms for the Mount Tobenna samples;
2. Eccentricity (long and short), Obliquity (long and short) terms for the Mount Raggeto samples.

Each of the above periodicities is attached to a specific stratal thickness and recurrent textures or lithofacies.

In forthcoming papers we shall discuss the application of wavelets to the pre-processing of the data, as well as the effect of other net architectures on the accuracy of the data.

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Appendix A

A.1. Neural net learning procedure

In this section we discuss in more detail the neural network learning procedure. First, the error back propagation algorithm (EBP) is presented and then we discuss the learning procedure, explaining the meaning of its main parameters. The algorithm is shown below. As already stated, it is necessary that the units have non-linear threshold functions that are continuously differentiable; such as, for instance, the sigmoidal function (Eq. (2)):

$$O_k = f \left[\sum_n w_{kn} O_n - B_k \right] \quad (6)$$

A.2. Algorithm

1. Let us assign small random values to weight and thresholds and a fixed small value to the error confidence;
2. input patterns $X_p = x_{p0}, \dots, x_{pn-1}$, where p is the index of the p th pattern and n is the number of the input nodes;
3. for the first hidden layer calculate the actual output:

$$O_{pj} = f \sum_{i=0}^{n-1} w_{ji} x_i - B_j$$

and pass it as input to the next layer and so on for the next layers. f is the activation function. At the output layer calculate the actual error by a comparison between the actual net output and relative corrected output;

4. after an entire cycle of input pattern presentations, calculate the whole error function: $E = 1/2 \times \sum_j (t_{pj} - O_{pj})^2$ on all input patterns p ;

5. adapt weights starting from the output layer and going backwards. Calculate: $W_{ji}(\text{new}) = W_{ji}(\text{old}) + \eta \times \delta_{pj} O_{pj} + \alpha \times (W_{ji}(\text{new}) - W_{ji}(\text{old}))$, where η and α are constants called learning rate and momentum, respectively, and: $\delta_{pj} = f'(O_{pj}) (t_{pj} - O_{pj})$ for output units, $\delta_{pj} = f'(O_{pj}) (\sum_k W_{kj} \times \delta_{pk})$ for hidden units, are the error terms for pattern p on node j ;
6. If $E > \text{error confidence}$ then go to step 2 otherwise stop.

The meaning of the hidden units in the net can be summarized as feature detectors. It can be seen as a recording of the inputs so that a neural net can learn the mapping of input patterns to output patterns. This recording, or internal representation, is critical to the behaviour of the network, in terms of the number of hidden units. In fact, to form internal representations of any input pattern, and to produce the correct response of output units, requires enough hidden units. The weight updating phase in the above algorithm, well known in literature as the 'generalized delta rule', provides a method for teaching multilayer perceptron networks, producing the necessary internal representation of the hidden nodes. It must be recalled that the learning procedure is not guaranteed to produce convergence. It is possible for the network to fall into a so-called *local minimum* of the error function E . The correct learning situation is, instead, the falling of the error function E to a so-called *absolute minimum*. In fact, the error function and its related network weight can be depicted as an energy surface that is a rippling landscape of hills and valleys, wells and mountains, with points of minimum energy corresponding to the wells and maximum energy found on the peaks [24].

In this context; the generalized delta rule aims to minimize the error function E by adjusting the weights so that they correspond to those at which the energy surface is the lowest. It does this by a method known as *gradient descent*, where the energy function is calculated and changes are made in the steepest downward direction. This is guaranteed to find a solution in cases where the energy landscape is simple. Each possible solution is represented as a hollow in the landscape. These 'basins of attractions' represent the solution to the values of the weights that produce the correct output. Three layers of neurons can, therefore, form arbitrarily complex

shapes, and are capable of separating any classes, as is stated by the well known Kolmogorov Theorem [24]. One of the main criticisms against the multilayer perceptron is that it requires many presentations of the input pattern set and a repetition of the corresponding calculations before the network is able to settle into a stable solution. The method of gradient descent is slow to converge, due to the complexity of the error surface. The addition of the two terms, learning rate and momentum, often speeds up convergence. The η term is a measure of the degree of influence of the updating weights formula of the error term, whereas the α term determines the influence of the past history of weight changes in the same formula. Determining the optimal choices for all these parameters must be achieved on the basis of a trial and error approach.

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